

LECTURE 35

The Planetary Boundary Layer (continued)

Ekman layer (continued)

Recap: we're solving the x - and y -component Reynolds-averaged equations of motion using an eddy viscosity closure for the Reynolds stresses. We combined two coupled ODEs for $\bar{u}_a \equiv \bar{u} - u_g$, and $\bar{v}_a \equiv \bar{v} - v_g$ into a single ODE for $M \equiv \bar{u}_a + i\bar{v}_a$,

$$0 = -ifM + K \frac{d^2M}{dz^2}. \quad (43)$$

This equation has solutions of the form $M \sim e^{qz}$ for two q :

$$q_1 = (1+i) \sqrt{\frac{f}{2K}}, \quad q_2 = -(1+i) \sqrt{\frac{f}{2K}}. \quad (44)$$

In terms of the parameter

$$\gamma \equiv \sqrt{\frac{f}{2K}}, \quad \text{which is real (in N. hemisphere) and positive,}$$

these q 's become: $q_1 = \gamma(1+i)$, $q_2 = -\gamma(1+i)$. The general solution of (43) is a linear combination of the solutions associated with these q :

$$M(z) = C \exp[\gamma(1+i)z] + D \exp[-\gamma(1+i)z], \quad (45)$$

where C and D are constants. Rewrite the first exponential:

$$\begin{aligned} \exp[\gamma(1+i)z] &= \exp(\gamma z + i\gamma z) = e^{\gamma z} e^{i\gamma z} && \text{use Euler's formula} \\ &= e^{\gamma z} (\cos \gamma z + i \sin \gamma z) \end{aligned}$$

Similarly, for the second exponential, we have

$$\begin{aligned}\exp[-\gamma(1+i)z] &= \exp(-\gamma z - i\gamma z) = e^{-\gamma z} e^{-i\gamma z} \\ &= e^{-\gamma z} (\cos \gamma z - i \sin \gamma z)\end{aligned}$$

Put 'em back into (45), get:

$$M(z) = C e^{\gamma z} (\cos \gamma z + i \sin \gamma z) + D e^{-\gamma z} (\cos \gamma z - i \sin \gamma z) \quad (46)$$

Now consider boundary conditions. Far above the surface ($z \rightarrow \infty$), the wind becomes geostrophic: $\bar{u} \rightarrow u_g$ and $\bar{v} \rightarrow v_g$. So $\bar{u}_a = \bar{u} - u_g \rightarrow u_g - u_g = 0$ and $\bar{v}_a \equiv \bar{v} - v_g \rightarrow v_g - v_g = 0$. So we impose:

$$\lim_{z \rightarrow \infty} M = \lim_{z \rightarrow \infty} (\bar{u}_a + i \bar{v}_a) = 0, \quad [\text{upper boundary condition for } M]$$

Looking at (46) as $z \rightarrow \infty$ we see that the term multiplying D goes to 0 since $e^{-\gamma z} \rightarrow 0$ (γ being positive), but the term multiplying C blows up since $e^{\gamma z} \rightarrow \infty$, (γ being positive). To ensure $M \rightarrow 0$, **we must set $C = 0$** . Then (46) becomes

$$M(z) = D e^{-\gamma z} (\cos \gamma z - i \sin \gamma z). \quad (47)$$

Now consider the surface conditions: no-slip on u and v ; $u(0) = 0$ and $v(0) = 0$. So $\bar{u}_a(0) = \bar{u}(0) - u_g = 0 - u_g = -u_g$ and $\bar{v}_a(0) = \bar{v}(0) - v_g = 0 - v_g = -v_g$. So:

$$M(0) = \bar{u}_a(0) + i \bar{v}_a(0) = -(\bar{u}_g + i \bar{v}_g). \quad [\text{lower boundary condition for } M]$$

So, set $z = 0$ in (47) and use the lower boundary condition for M :

$$M(0) = D e^{-0} [\cos(0) - i \sin(\gamma 0)] \quad \text{plug in lower b.c. for } M$$

$$-(\bar{u}_g + i \bar{v}_g) = D [1 - 0] = D \quad \therefore \boxed{D = -(\bar{u}_g + i \bar{v}_g)}$$

So now (47) becomes

$$M(z) = -(\bar{u}_g + i\bar{v}_g)e^{-\gamma z}(\cos\gamma z - i\sin\gamma z).$$

Plug in $M \equiv \bar{u}_a + i\bar{v}_a$.

$$\bar{u}_a + i\bar{v}_a = -(\bar{u}_g + i\bar{v}_g)e^{-\gamma z}(\cos\gamma z - i\sin\gamma z) \quad (48)$$

To find out what \bar{u}_a and \bar{v}_a are, we need to take real and imaginary parts of (48). So, rewrite the right hand side of (48) to make clear what its real and imaginary parts are:

$$\begin{aligned} \bar{u}_a + i\bar{v}_a &= -e^{-\gamma z} \left[(\bar{u}_g + i\bar{v}_g) (\cos\gamma z - i\sin\gamma z) \right] \quad \text{expand out right hand side} \\ &= -e^{-\gamma z} \left[\bar{u}_g \cos\gamma z - i\bar{u}_g \sin\gamma z + i\bar{v}_g \cos\gamma z - i(i\bar{v}_g)\sin\gamma z \right] \end{aligned}$$

In the last term use fact that $-ii = -i^2 = -(-1) = 1$. So we get

$$\bar{u}_a + i\bar{v}_a = -e^{-\gamma z} \left(\bar{u}_g \cos\gamma z + \bar{v}_g \sin\gamma z - i\bar{u}_g \sin\gamma z + i\bar{v}_g \cos\gamma z \right) \quad (49)$$

Now take real and imaginary parts of this equation. The real part gives:

$$\bar{u}_a = -e^{-\gamma z} \left(\bar{u}_g \cos\gamma z + \bar{v}_g \sin\gamma z \right)$$

The imaginary part gives:

$$\bar{v}_a = -e^{-\gamma z} \left(-\bar{u}_g \sin\gamma z + \bar{v}_g \cos\gamma z \right)$$

Lastly, rewrite these in terms of the mean winds $\bar{u} = u_g + \bar{u}_a$, and $\bar{v} = v_g + \bar{v}_a$:

$$\bar{u} = \bar{u}_g \left(1 - e^{-\gamma z} \cos\gamma z \right) - \bar{v}_g e^{-\gamma z} \sin\gamma z, \quad (50)$$

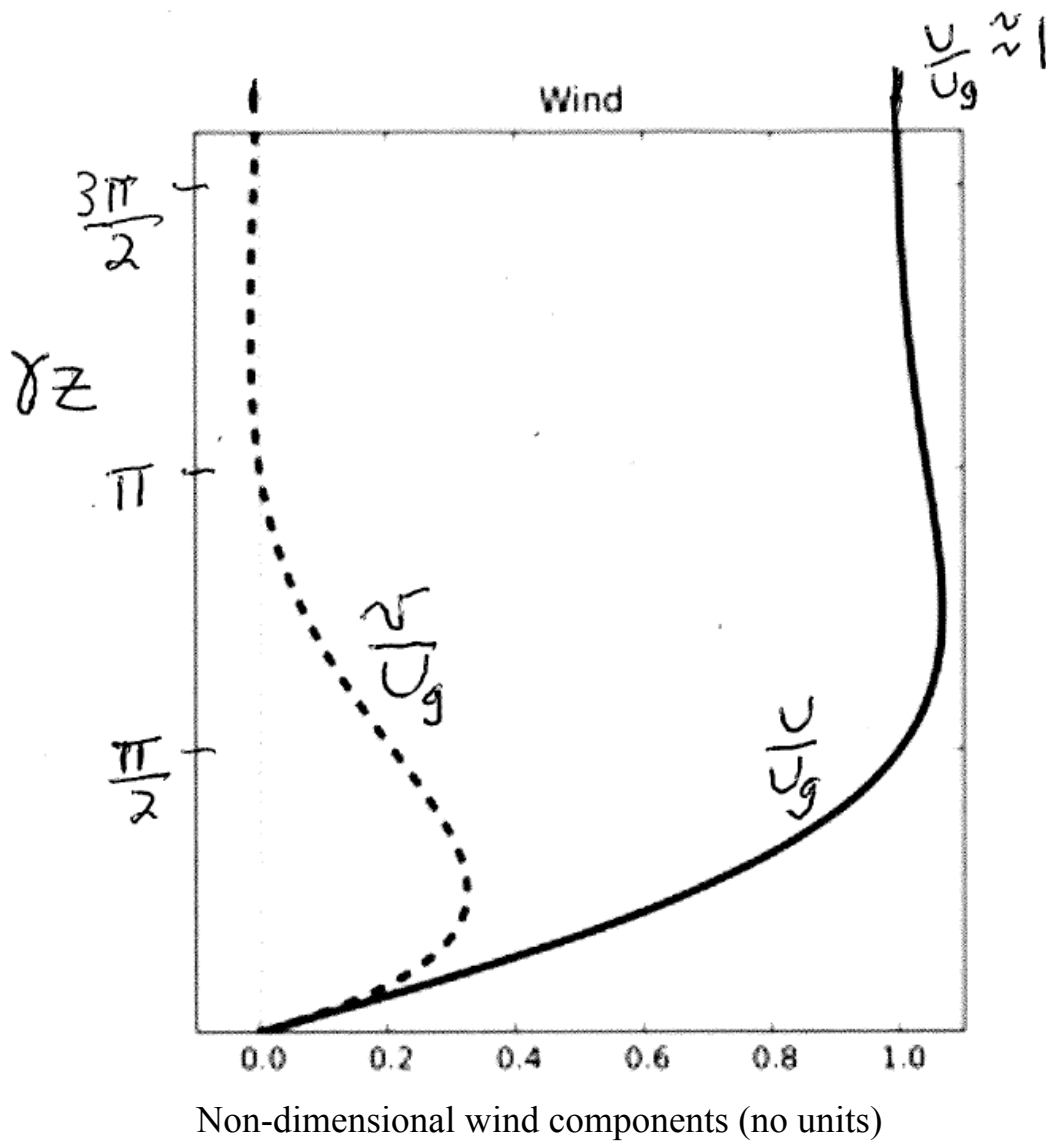
$$\bar{v} = \bar{v}_g \left(1 - e^{-\gamma z} \cos\gamma z \right) + \bar{u}_g e^{-\gamma z} \sin\gamma z, \quad (51)$$

where $\gamma \equiv \sqrt{\frac{f}{2K}}$.

[Reality check: Verify using (50) and (51) that flow becomes geostrophic far above surface and goes to zero at surface. Yep, it checks out.]

Consider a new (rotated) coordinate system in which **new x axis is aligned with geostrophic wind** (all of geostrophic wind vector is in new x-direction, so $v_g = 0$). So **y axis points toward low pressure!**). In this new system, (50) and (51) become:

$\bar{u} = \bar{u}_g (1 - e^{-\gamma z} \cos \gamma z)$, $\bar{v} = \bar{u}_g e^{-\gamma z} \sin \gamma z$. These are graphed below:



Within boundary layer there's a **cross-isobar flow toward low pressure** ($v > 0$).

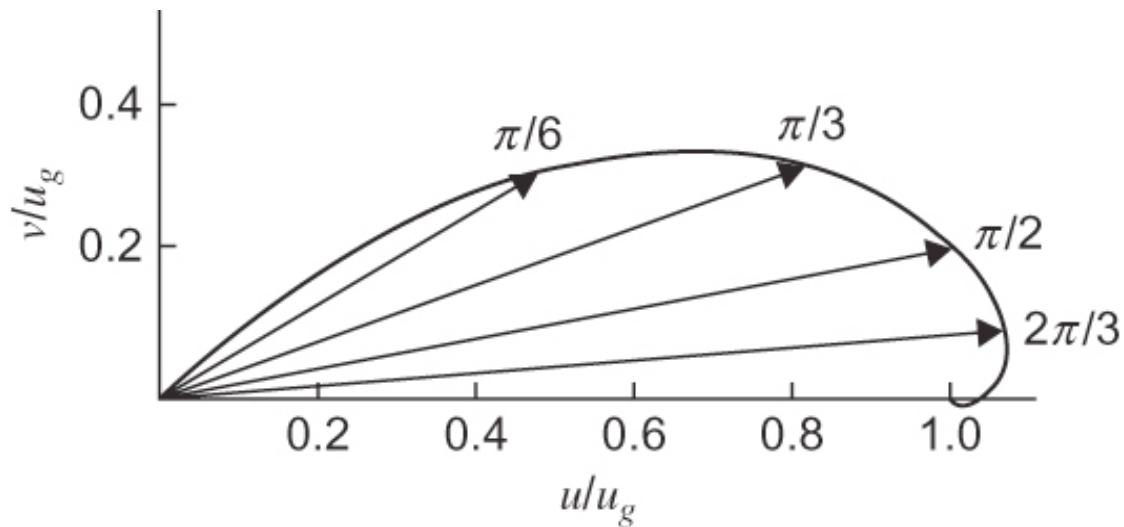
Various ways to define height of boundary layer. One way is to say it's the height H where $\gamma H = \pi$ (so $e^{-\gamma H} = e^{-\pi} \approx 0.05$). At this height u has nearly reached its geostrophic value and v has nearly returned to 0 (see diagram above).

$$H = \frac{\pi}{\gamma} = \frac{\pi}{\sqrt{\frac{f}{2K}}} = \pi \sqrt{\frac{2K}{f}}$$

For typical values $K = 10 \text{ m}^2 \text{ s}^{-1}$, $f = 8 \times 10^{-5} \text{ s}^{-1}$, we obtain H as: $H \sim 1.5 \text{ km}$.

Ekman spiral

(hodograph of v versus u for γz ranging from ground through top of boundary layer):



Theoretical Ekman spiral (top dashed curve) versus a real Ekman spiral (solid curve):

