

Solutions to Problem Set #4
METR 5413
Advanced Synoptic Meteorology
H. Bluestein

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(a)

$$\theta = T \left(\frac{p_0}{p} \right)^\kappa, \text{ where } \kappa = \frac{R}{c_p} = \frac{287}{1004} = 0.286$$

$$\theta_{850} = (284 \text{ K}) \left(\frac{100 \text{ kPa}}{85 \text{ kPa}} \right)^{0.286} = 297.5 \text{ K}$$

$$\theta_{500} = (255 \text{ K}) \left(\frac{100 \text{ kPa}}{50 \text{ kPa}} \right)^{0.286} = 310.9 \text{ K}$$

$$-\frac{\partial \theta}{\partial p} = -\frac{297.5 \text{ K} - 310.9 \text{ K}}{85 \text{ kPa} - 50 \text{ kPa}} = 0.383 \text{ K kPa}^{-1}$$

$$\begin{aligned} \theta_{700} &= \theta_{500} + (70 \text{ kPa} - 50 \text{ kPa}) \frac{\partial \theta}{\partial p} \\ &= 310.9 \text{ K} + (20 \text{ kPa})(-0.383 \text{ K kPa}^{-1}) \\ &= 303.2 \text{ K} \end{aligned}$$

$$T_{700} = (303.2 \text{ K}) \left(\frac{70 \text{ kPa}}{100 \text{ kPa}} \right)^{0.286} = 273.8 \text{ K}$$

$$\begin{aligned} \sigma &= -\frac{RT}{p\theta} \frac{\partial \theta}{\partial p} \\ &= \frac{(287 \text{ m}^2 \text{ s}^{-2} \text{ K}^{-1})(273.8 \text{ K})(0.383 \text{ K kPa}^{-1})}{(70 \text{ kPa})(303.2 \text{ K})} \\ &= 1.42 \text{ m}^2 \text{ s}^{-2} \text{ kPa}^{-2} \end{aligned}$$

(b)

$$\sigma_e = \sigma - \frac{0.622 R L_v}{p^2 c_p} \left[\frac{de_s}{dT} \left(\frac{dT}{dp} \right)_m - \frac{e_s}{p} \right]$$

$$e_s = 5.9728 \times 10^{22} e^{-6788.026/T} T^{-5.0065} \text{ kPa}$$

$$\frac{de_s}{dT} = e_s \left[\frac{6788.026}{T^2} - \frac{5.0065}{T} \right] \text{ kPa K}^{-1}$$

$$\left(\frac{dT}{dp} \right)_m = \frac{RT}{pc_p} \left[\frac{1 + \frac{L_v r_s}{RT}}{1 + \left(\frac{0.622 L_v}{c_p T} \right) \left(\frac{L_v r_s}{RT} \right)} \right]$$

$$c_s(273.8 \text{ K}) = 0.64 \text{ kPa}$$

$$\frac{dc_s}{dT}(273.8 \text{ K}) = 0.0462 \text{ kPa K}^{-1}$$

$$r_s = 0.622 \frac{c_s}{p} = 0.622 \frac{0.64 \text{ kPa}}{70 \text{ kPa}} = 5.69 \times 10^{-3}$$

$$\left(\frac{dT}{dp}\right)_m = 0.654 \text{ K kPa}^{-1}$$

$$\begin{aligned} \sigma_e &= \sigma - \frac{(0.622)(287 \text{ m}^2 \text{ s}^{-2} \text{ K}^{-1})(2.5 \times 10^6 \text{ m}^2 \text{ s}^{-2})}{(70 \text{ kPa})^2 (1004 \text{ m}^2 \text{ s}^{-2} \text{ K}^{-1})} \left[(0.0462 \text{ kPa K}^{-1})(0.654 \text{ K kPa}^{-1}) - \frac{0.64 \text{ kPa}}{70 \text{ kPa}} \right] \\ &= 1.42 \text{ m}^2 \text{ s}^{-2} \text{ kPa}^{-2} - 1.91 \text{ m}^2 \text{ s}^{-2} \text{ kPa}^{-2} \\ &= -0.49 \text{ m}^2 \text{ s}^{-2} \text{ kPa}^{-2} \end{aligned}$$

Therefore, there is conditional instability, and the operator in the Q-G ω equation is not elliptic. ("Moist convective adjustment" would occur in numerical models.)

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parabolic: $\omega = ap^2 + bp + c$, $\frac{\partial \omega}{\partial p} = 2ap + b$

$$\omega(p=100 \text{ kPa}) = 0 = a(100 \text{ kPa})^2 + b(100 \text{ kPa}) + c$$

$$\omega(p=60 \text{ kPa}) = -10 \mu\text{b s}^{-1} = a(60 \text{ kPa})^2 + b(60 \text{ kPa}) + c$$

$$\frac{\partial \omega}{\partial p}(p=60 \text{ kPa}) = 0 = 2a(60 \text{ kPa}) + b$$

solution for a, b, and c:

$$a = 6.25 \times 10^{-7} \text{ kPa}^{-1} \text{ s}^{-1}, \quad b = -7.5 \times 10^{-5} \text{ s}^{-1}, \quad c = 1.25 \times 10^{-3} \text{ kPa s}^{-1}$$

$$\begin{aligned} \delta(p=800 \text{ mb}) &= -\frac{\partial \omega}{\partial p}(p=800 \text{ mb}) \\ &= -[2(6.25 \times 10^{-7} \text{ kPa}^{-1} \text{ s}^{-1})(80 \text{ kPa}) - 7.5 \times 10^{-5} \text{ s}^{-1}] \\ &= -2.5 \times 10^{-5} \text{ s}^{-1} \end{aligned}$$

$$\frac{\partial}{\partial t} \left(-\frac{\partial \theta}{\partial p} \right) = -\delta \frac{\partial \theta}{\partial p}$$

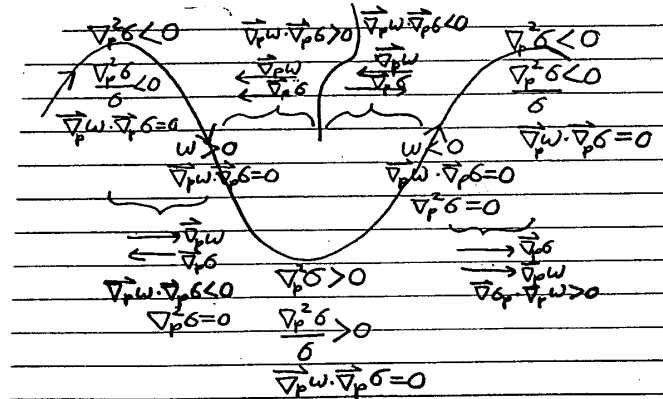
$$\frac{d \left(-\frac{\partial \theta}{\partial p} \right)}{-\frac{\partial \theta}{\partial p}} = \delta \, dt$$

$$\begin{aligned} \ln \left(-\frac{\partial \theta}{\partial p} \right) \Big|_0^t &= \delta t \\ &= \ln \left(\frac{0.001 \text{ K kPa}^{-1}}{0.383 \text{ K kPa}^{-1}} \right) \\ &= -5.95 \end{aligned}$$

therefore,

$$t = \frac{-5.95}{\delta(p=800 \text{ mb})} = \frac{-5.95}{-2.5 \times 10^{-5} \text{ s}^{-1}} = 2.38 \times 10^5 \text{ s} = 2.76 \text{ days}$$

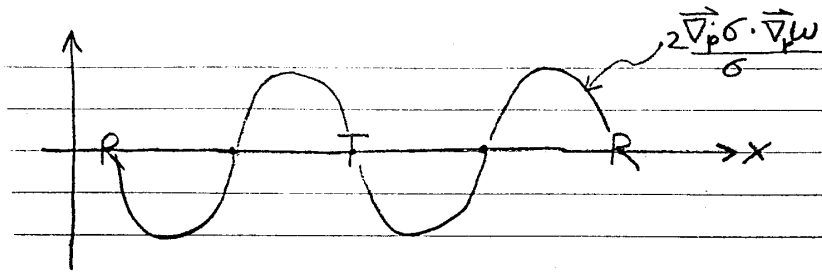
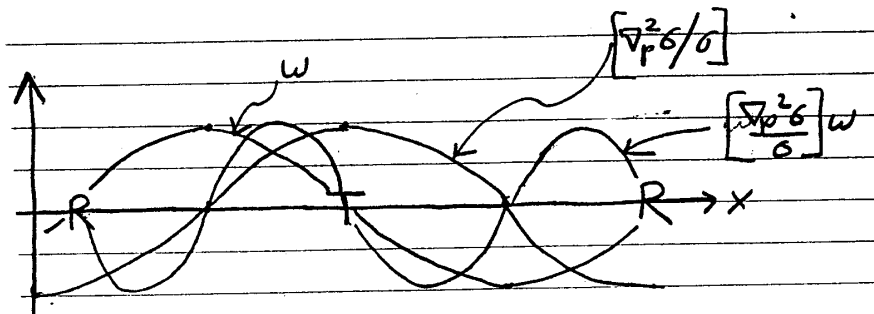
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The effect of horizontal variations in σ on vertical motion is contained in the terms

$$\frac{\nabla_p^2 \sigma}{\sigma} w \text{ and } \frac{2}{\sigma} \nabla_p \sigma \cdot \nabla_p w.$$

The following diagrams show the general forms of these quantities.



Therefore, effects of horizontal variations in σ are *least* important along trough and ridge axes, and exactly in between trough and ridge axes. Effects of horizontal variations in σ are *most* important $1/8$ wavelength downstream and upstream from trough and ridge axes.