Answers to Homework Assignment #1

1.1 $r_a = cT_s/2$ and $v_a = \lambda/4T_s$

where c is the velocity of the microwaves, λ is its wavelength $\lambda = c / f$, *f* is the wave's frequency, and T_s is the pulse repetition time (PRT; i.e., the interval between echo samples)

- 1.2 The observed Doppler velocity v_0 equals the true Doppler velocity v_t if v_t lies in the unambiguous interval $\pm v_a$. If v_t exceeds $+v_a$, the observed velocity v_0 aliases to negative velocities. Thus as v_t increases beyond $+v_a$, the negative values of v_0 decrease as $v_t 2v_a$.
- 1.3 First trip ground clutter from scatterers on the ground at ranges, for example, between zero and 20 km is overlaid on 2^{nd} trip weather echoes if weather scatterers are at a range r_a plus 0 to 20 km.
- 1.4 (a) Let the range error be denoted as Δr . Then

$$\Delta r = ct - vt = ct - \frac{c}{n}t = ct\left(1 - \frac{1}{n}\right)$$

where *t* is the time after the microwave pulse is transmitted. But for a scatter at r = 300 km, t = r/v = rn/c. Therefore $\Delta r = r(n-1)$. Given n = 1.000300, and r = 300 km,

$$\Delta r = 3 \times (3 \times 10^{-4}) = 90 \,\mathrm{m}$$

(b) The echo phase difference $\Delta \psi_{e}$ is

$$\Delta \psi_{e} = -2\pi f \left(\frac{2r}{c} - \frac{2r}{v}\right) = -4\pi f \frac{r}{c}(1-n) = 4\pi \frac{r}{\lambda}(n-1)$$
$$= \frac{4\pi}{0.1} \times 3 \times 10^{4} (3 \times 10^{-4}) = 360\pi \text{ (radians)}$$

1.5 (a) P_w is the water vapor pressure:

$$P_w = 6.11(10)^{7.5 \left(\frac{T}{T+237.5}\right)}$$

N is the exponential reference profile of the refractivity:

$$N = 313 \ e^{-0.1439h}$$

 N_{d} is the refractivity profile computed from the sounding neglecting the water vapor effect:

$$N_d = \frac{77.6P}{T + 273.15}$$

 N_w is the refractivity which has taken into account the water vapor effect:



(b) The departure of N_d from the reference profile (N) may be attributed to the temperature and water vapor variations in the real atmosphere.

The difference between N_d and N_w is caused by the water vapor effect. In the lower part of the atmosphere where moisture is abundant, the difference between N_d and N_w is the largest.

(c) Use Fig. 1 to calculate $\frac{dN}{dh}$ near the surface.

For the moist atmosphere:

$$\frac{dN}{dh} \approx -7.64(10)^{-2}(m^{-1})$$

$$\frac{dn}{dh} \approx -7.64(10)^{-5}(km^{-1})$$

The corresponding effective earth's radius is:

$$a_e = \frac{a}{1 + a\left(\frac{dn}{dh}\right)} = \frac{6370}{1 + 6370[-7.64(10^{-5})]}$$
$$a_e = 12409 \ km$$

(d) Again, use Fig. 1 to calculate $\frac{dN}{dh}$ near the surface.

For exponential atmosphere:

$$\frac{dN}{dh} \approx -4.11(10)^{-2}(m^{-1})$$
$$\frac{dn}{dh} \approx -4.11(10)^{-5}(km^{-1})$$

The corresponding effective earth's radius is:

$$a_e = \frac{6370}{1 + 6370[-4.11(10^{-5})]}$$
$$a_e = 8629 \ km$$

For the dry atmosphere:

$$\frac{dN}{dh} \approx -3.30(10)^{-2}(m^{-1})$$
$$\frac{dn}{dh} \approx -3.30(10)^{-5}(km^{-1})$$

The corresponding effective earth's radius is:

$$a_e = \frac{6370}{1 + 6370[-3.30(10^{-5})]}$$
$$a_e = 8065 \ km$$

(e) The a_e for the exponential reference atmosphere agrees well with $\frac{4}{3}a$. But the a_e obtained from the actual sounding is much larger than the a_e for the $\frac{4}{3}$ earth radius model.



$$h \cong \frac{D^2}{2a_e}$$

For the observed atmosphere:

$$h = \frac{10^4}{2(12409)} km = 403 m$$

For the $\frac{4}{3}$ earth radius model:

$$h = \frac{10^4}{2\left(\frac{4}{3}\right)6370} km = 589 m$$

Therefore,

$$\Delta h \cong 200 m$$