

# **DOPPLER SPECTRA OF WEATHER SIGNALS**

(Chapter 5; examples from chapter 9)

# Spectrum and Autocorrelation

$$S(f) = \lim(M \rightarrow \infty) T_s \sum_{l=-(M-1)}^{M-1} R(l) e^{-j2\pi f T_s l} \quad (5.18)$$

$$R(l) = \int_{-1/2T_s}^{1/2T_s} S(f) e^{j2\pi f T_s l} df \quad \text{Skip?} \quad (5.19)$$

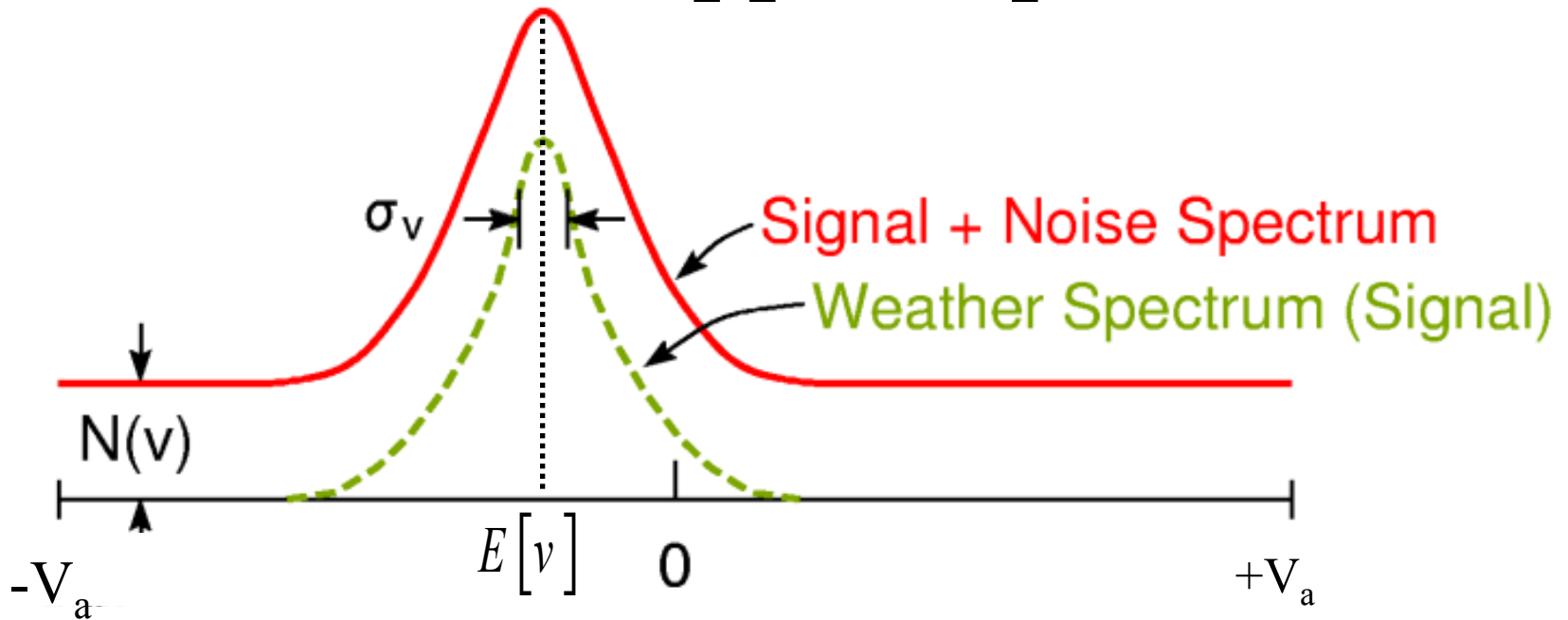
Two methods to calculate the power spectrum :

(1)  $\hat{S}(f)$  from the D. F. T. of  $\hat{R}(l)$  (5.21)

or

(2)  $\hat{S}(f)$  from  $|\text{F. T. of } V(n)|^2$  (5.22)

# Idealized Doppler Spectrum

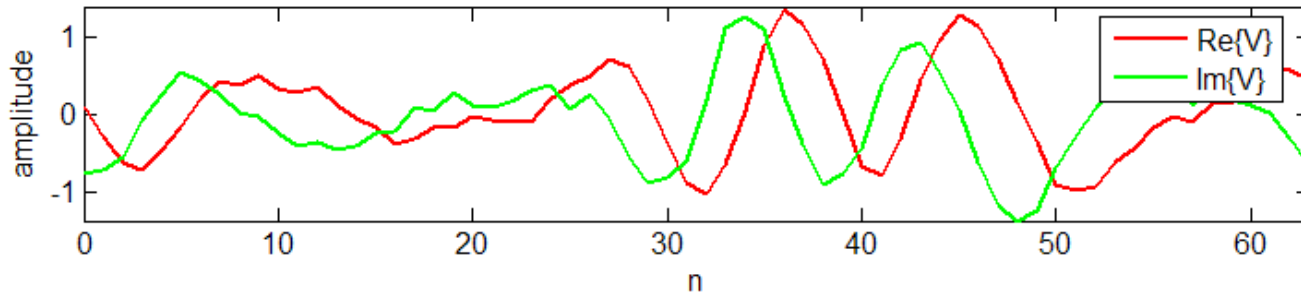


$$S(\nu) = \frac{S}{\sigma_\nu \sqrt{2\pi}} e^{-(\nu - \bar{\nu})^2 / 2\sigma_\nu^2} + \frac{N}{\lambda / 2T_s} \quad (6.3)$$

$$R(mT_s) = S e^{-8(\pi\sigma_\nu mT_s / \lambda)^2 + j4\pi\bar{\nu}mT_s / \lambda} + N\delta_m \quad (6.4)$$

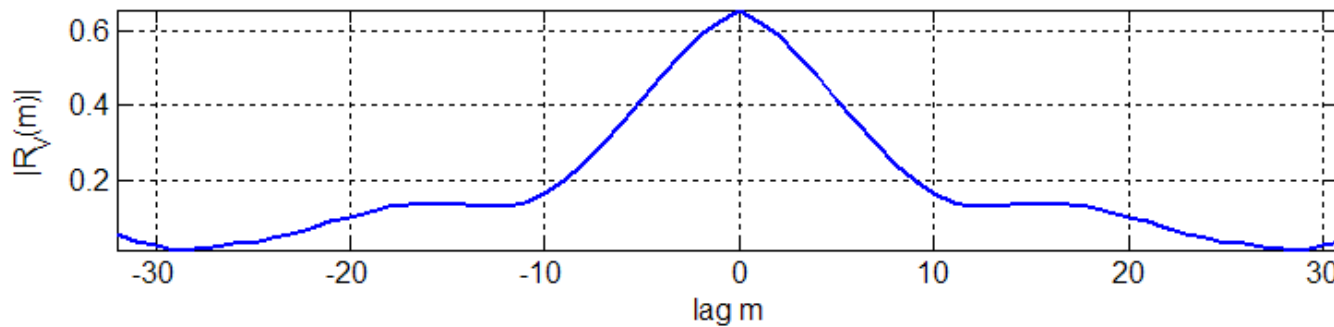
# Weather Signal Analysis

3 Domains:

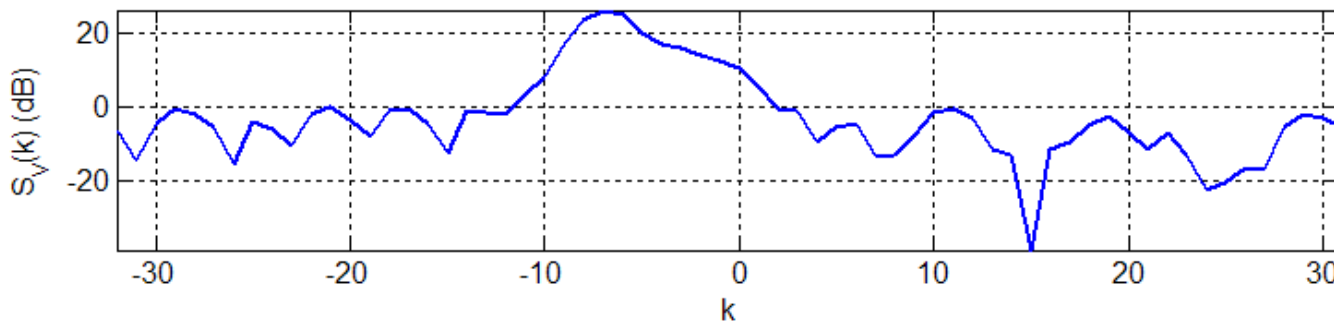


Time

$$|R_v(m)| = \frac{1}{M} \left| \sum_{n=0}^{|m|-1} V^*(n)V(n+m) \right| \quad (5.16)$$



Time Lag



Frequency

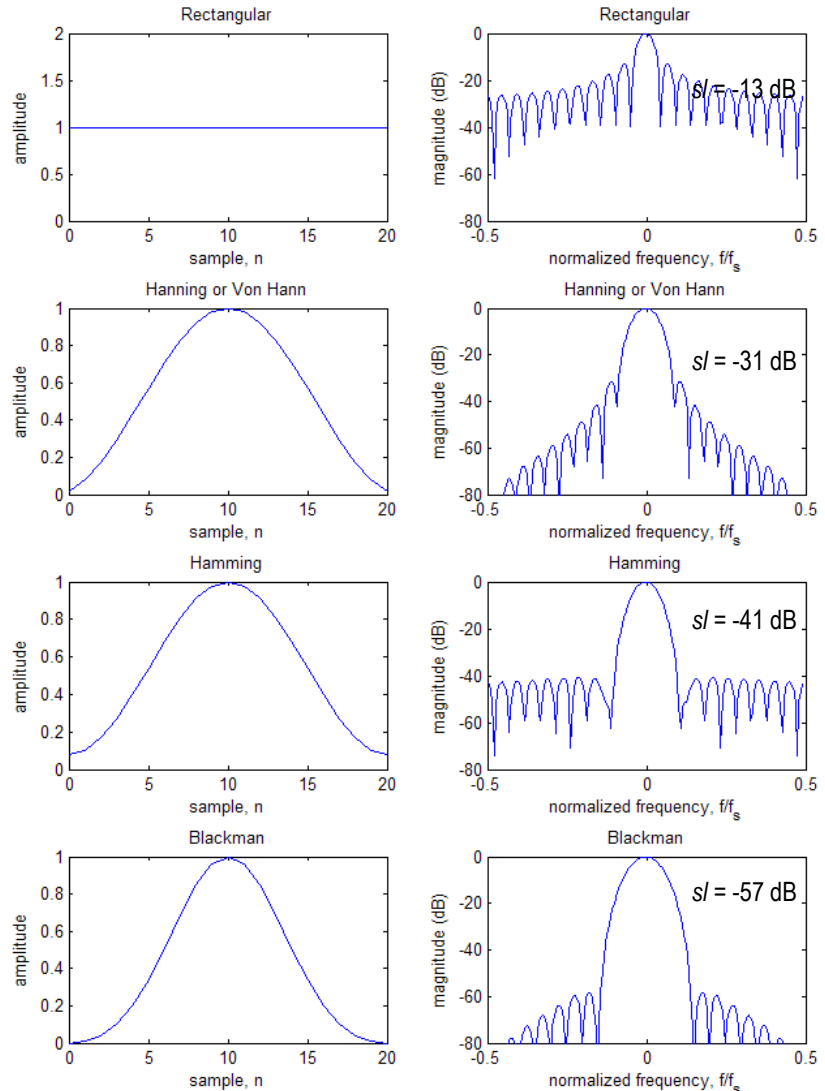
# Data Windowing

- Windowing forces the amplitude of the digital signal at both ends to go smoothly towards zero
  - Windows with lower frequency sidelobes reduce spectral leakage
  - Windows with wider frequency mainlobe reduce the frequency resolution

## • Windowed DFT

$$Z(k) = \sum_{m=0}^{M-1} d(m)V(m)e^{-j\frac{2\pi mk}{M}},$$

where  $d(m)$  is the window function  
and  $k = 0, 1, \dots, M - 1$



# Spectral Window Effect

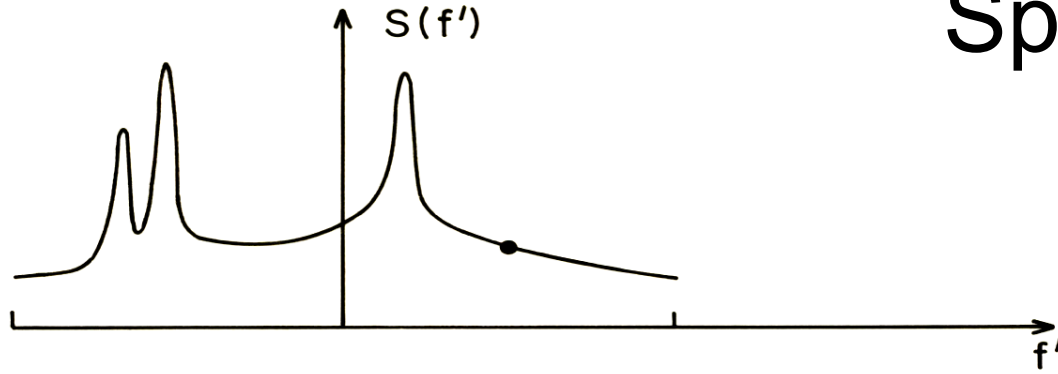
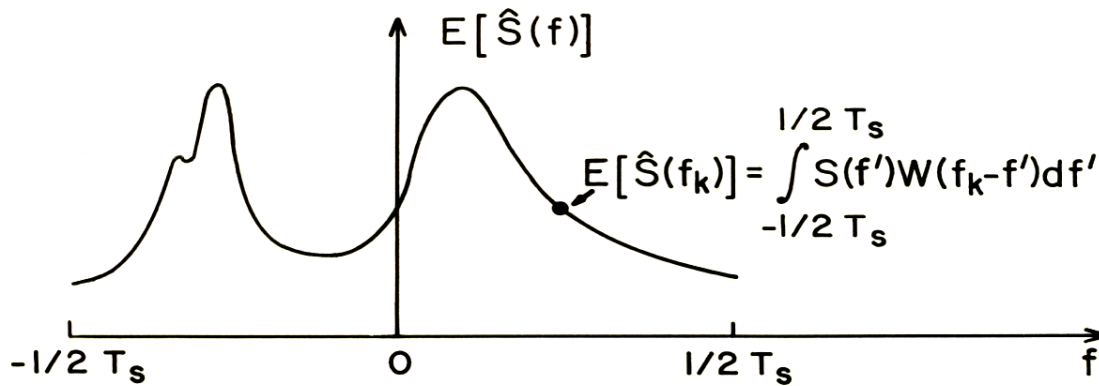
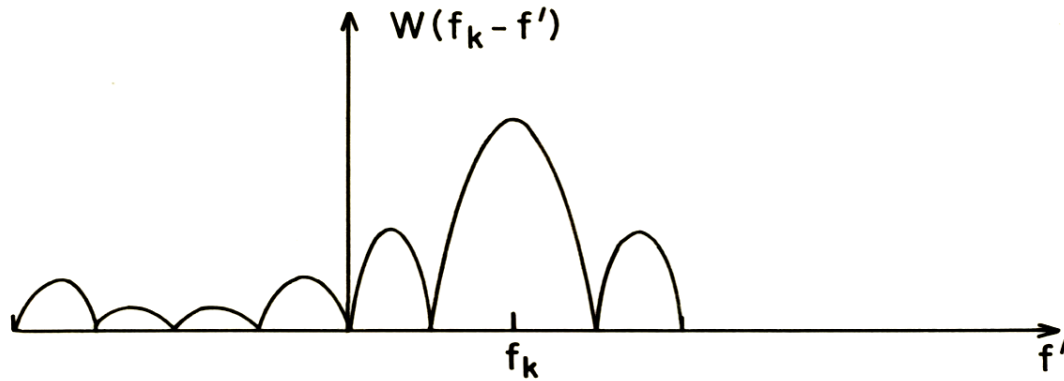
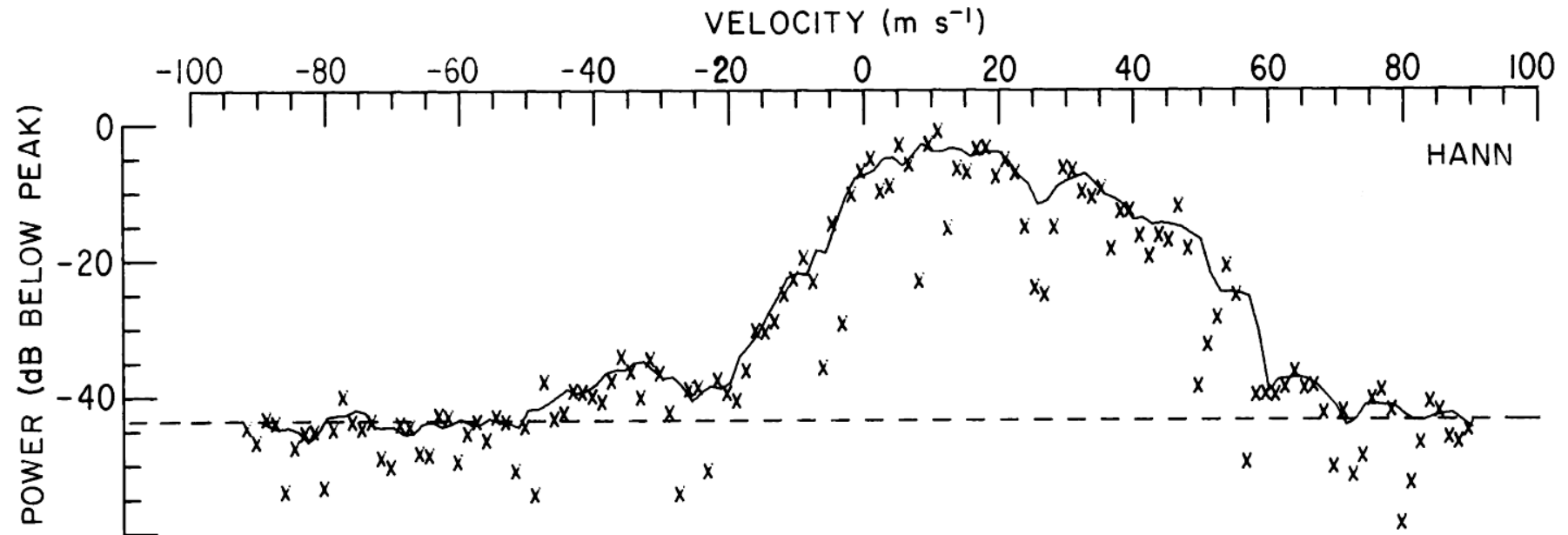
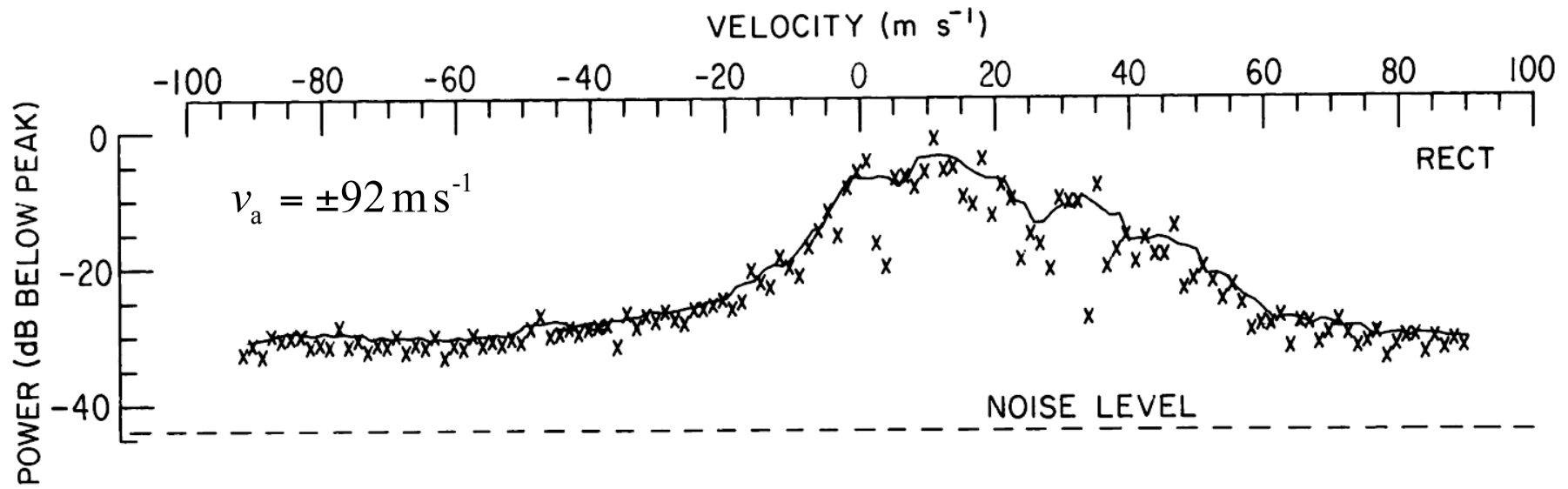


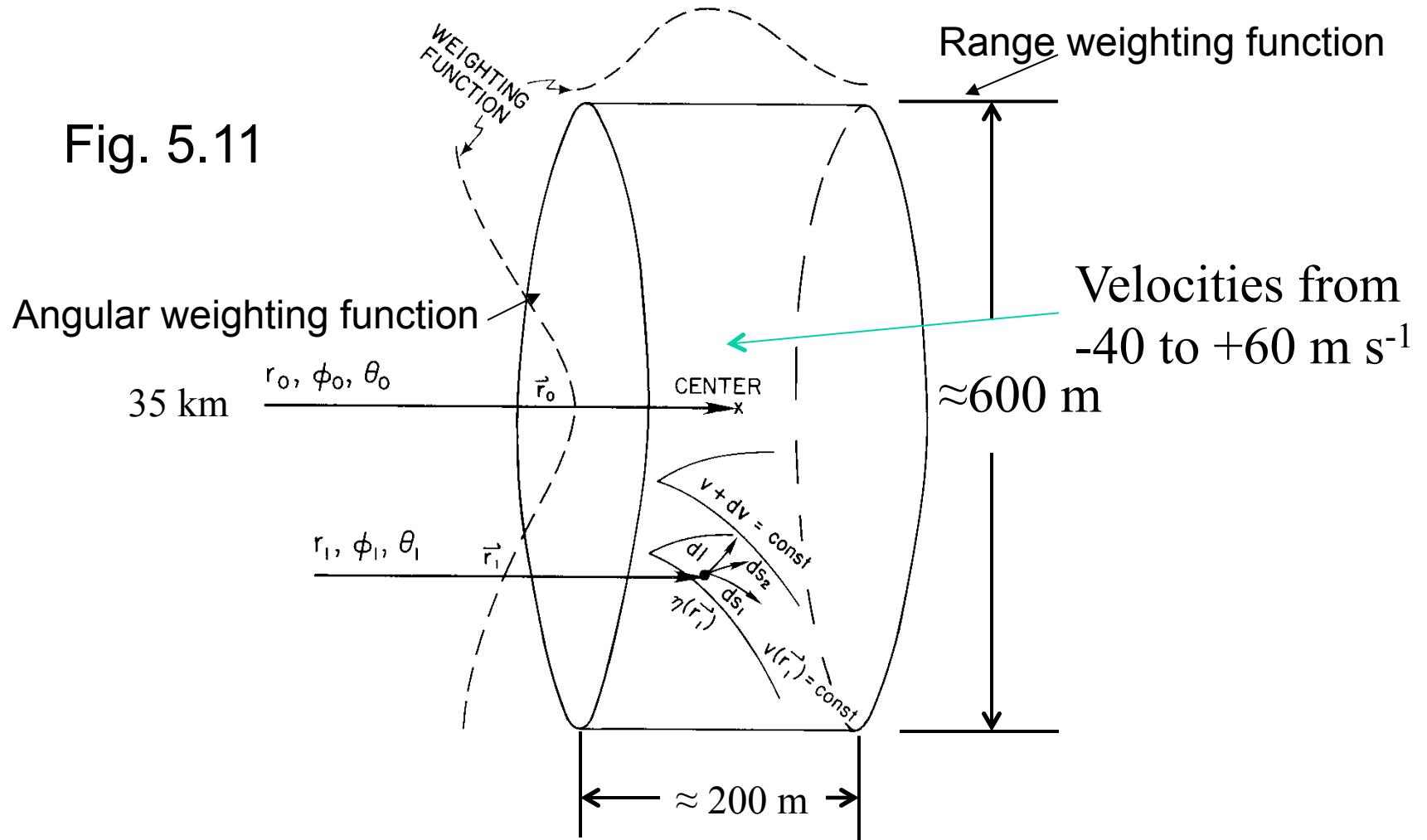
Fig. 5.8



# Window Effect on Spectra (Fig. 5.10)



# The Resolution Volume $V_6$





The Expected Doppler Velocity  
(i.e., the first moment of the Doppler spectrum)  
expressed as a Weighted spatial average

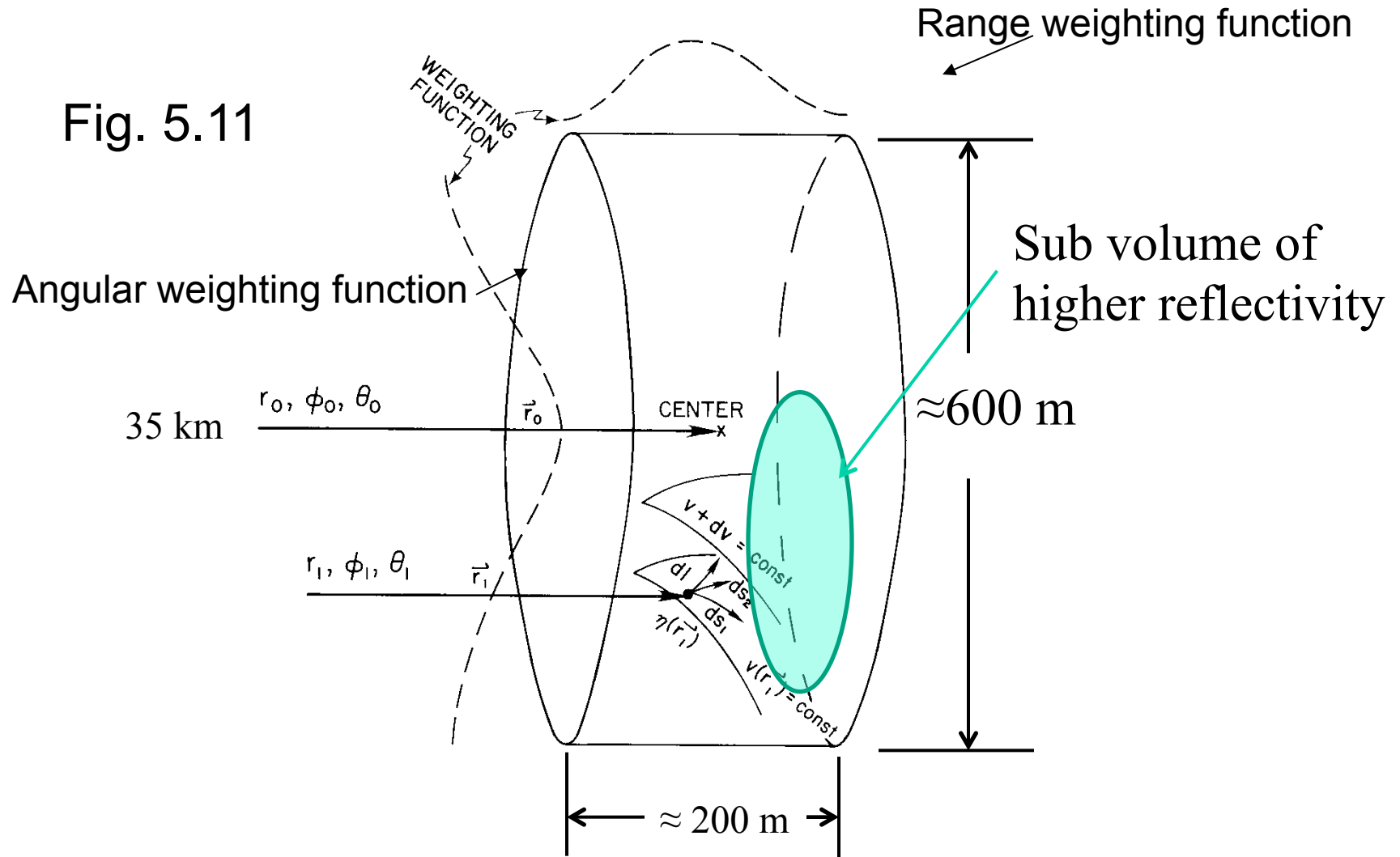
$$E[\bar{v}] = \frac{\iiint v(\mathbf{r}_1) \eta(\mathbf{r}_1) I(\mathbf{r}_0, \mathbf{r}_1) dV_1}{\iiint \eta(\mathbf{r}_1) I(\mathbf{r}_0, \mathbf{r}_1) dV_1} \quad (5.48)$$

$$I(\mathbf{r}_0, \mathbf{r}_1) = \frac{Cf^4 (\theta_1 - \theta_0, \phi_1 - \phi_0) |W_s(r_0, r_1)|^2}{l^2(\mathbf{r}_1) r_1^4} \quad (5.40)$$

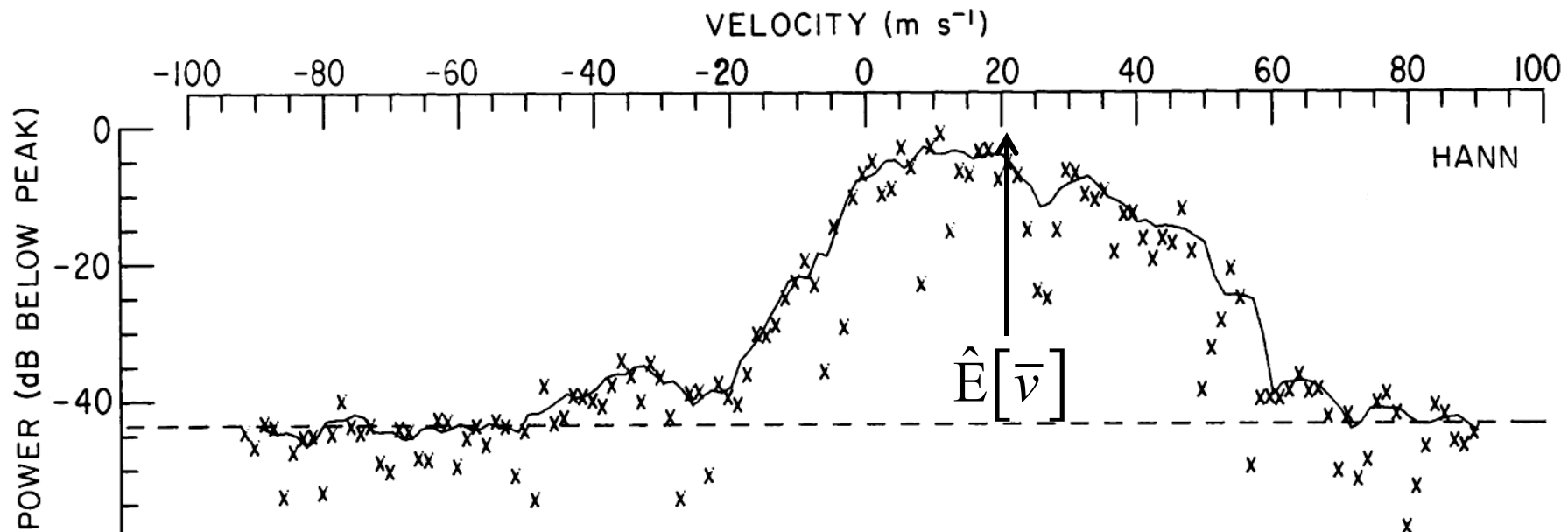
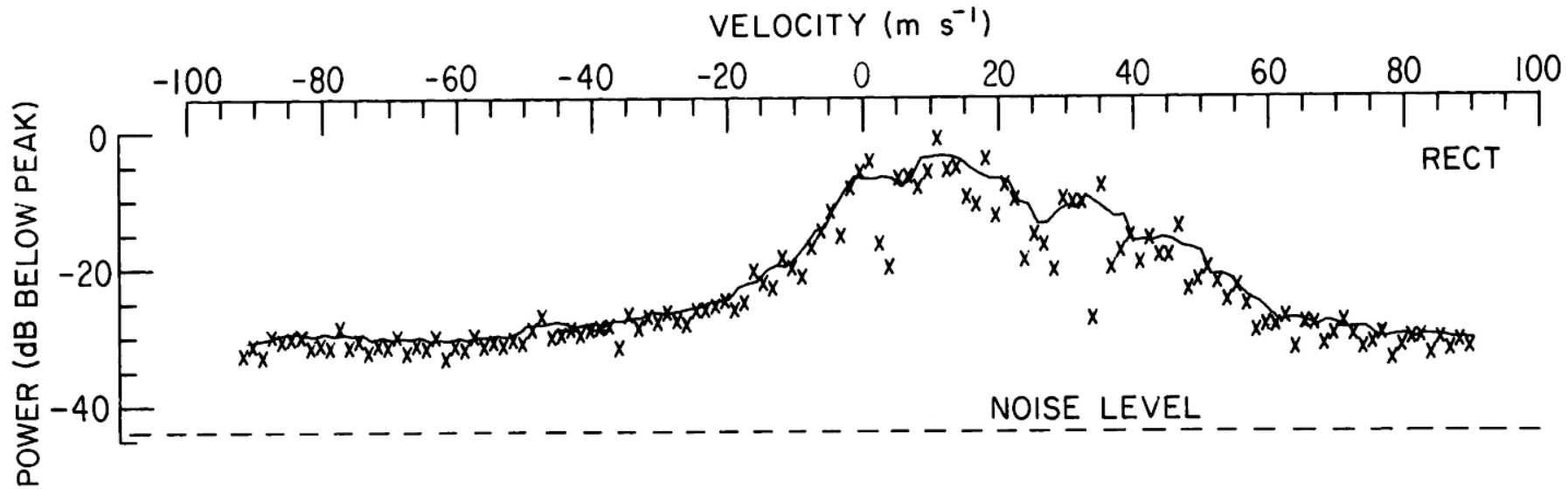
$I(\mathbf{r}_0, \mathbf{r}_1)$  is the Weighting Function

# The Resolution Volume $V_6$

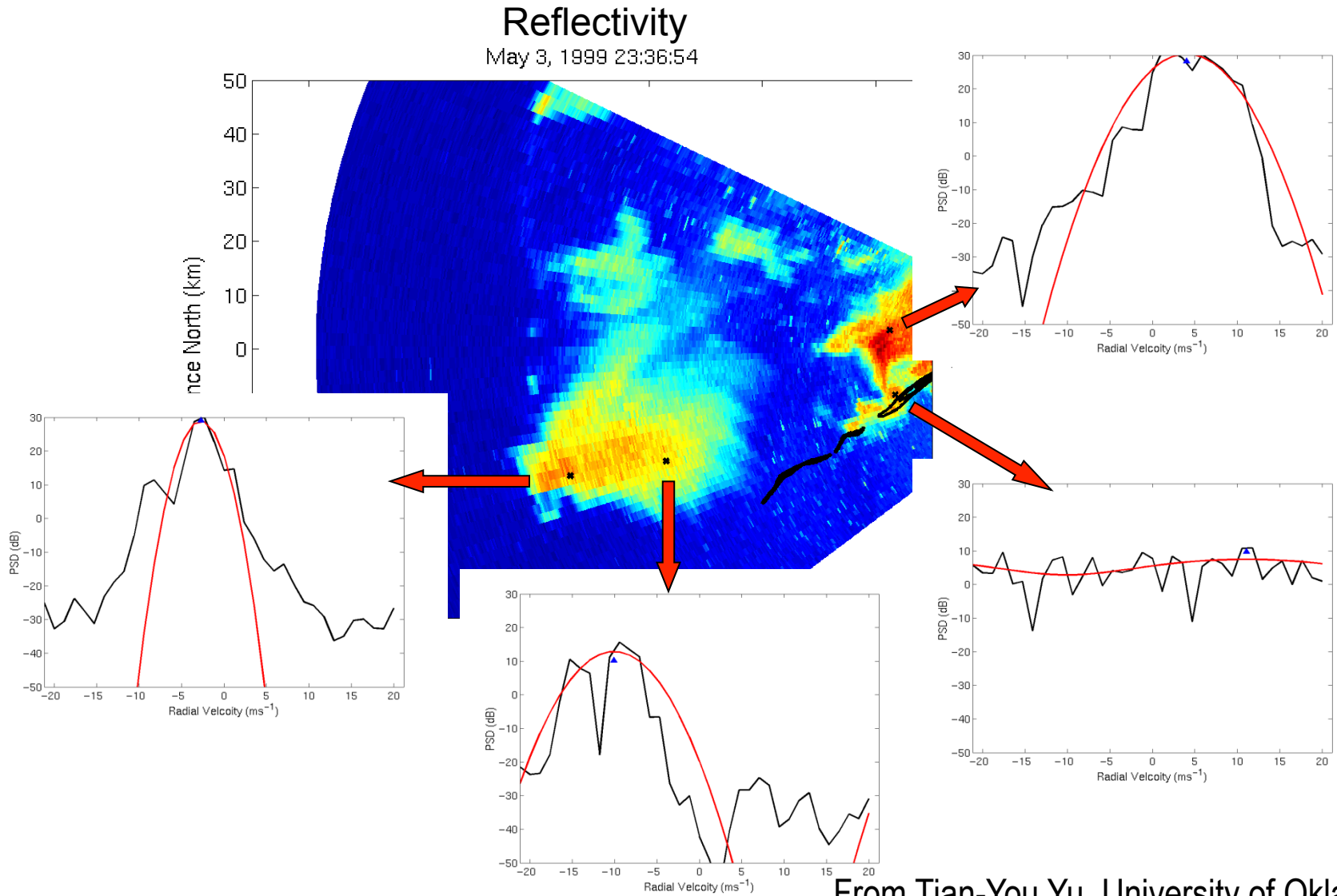
(typically not uniformly filled with scatterers)



# Window Effect on Spectra (Fig. 5.10)



# Examples of Doppler Spectra

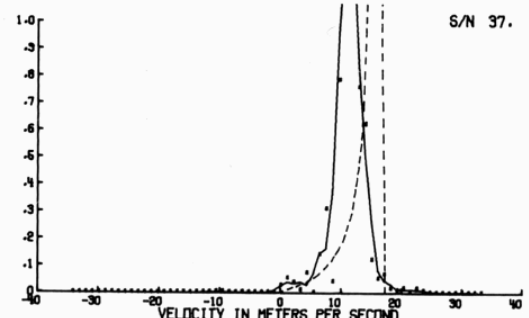
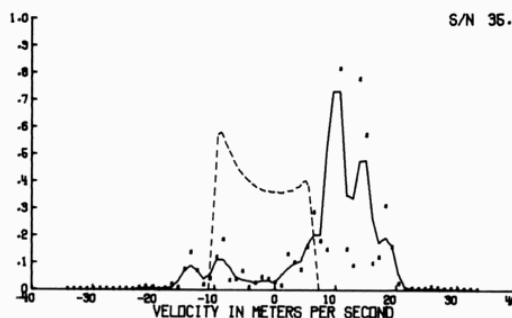
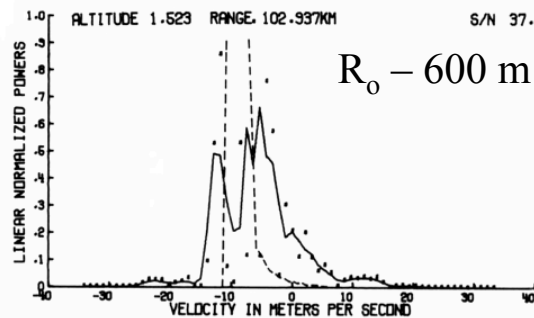
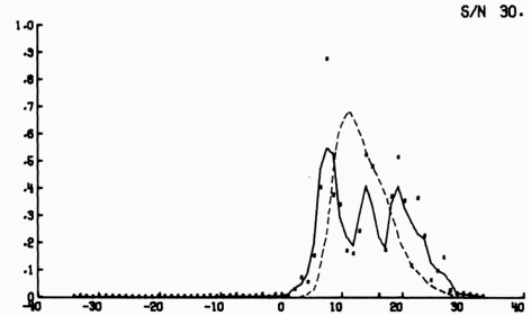
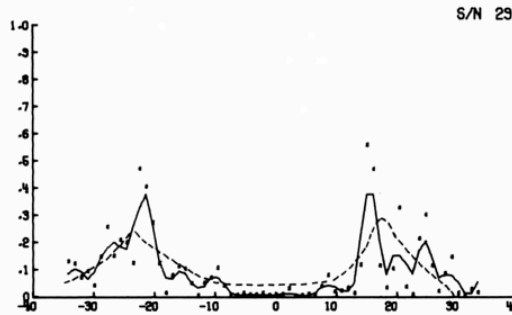
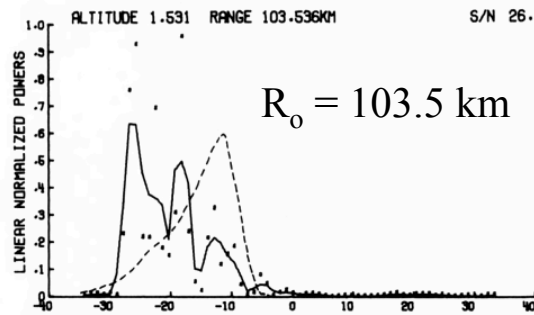
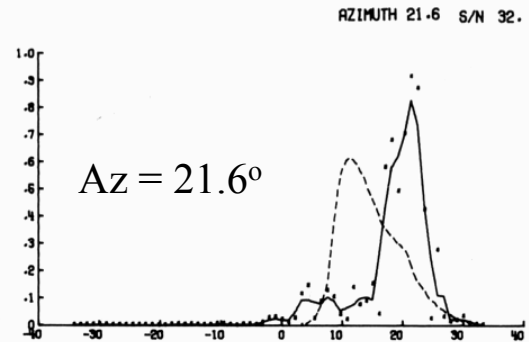
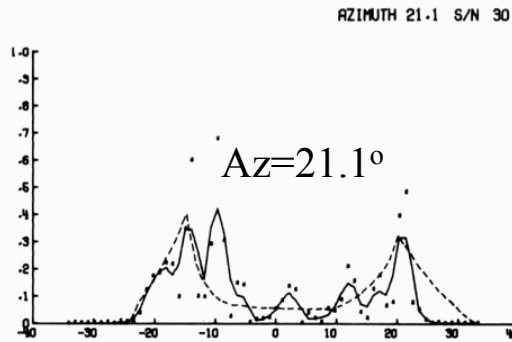
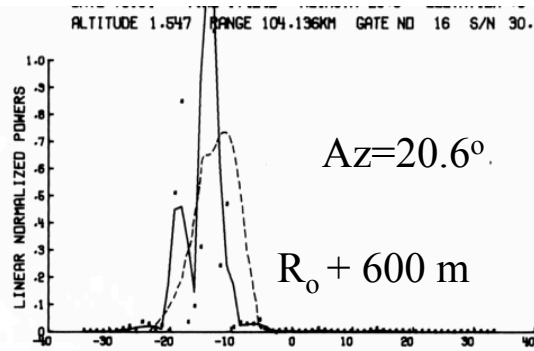


# Fitted and Observed Spectra in Resolution Volumes Surrounding the Stillwater Tornado (Similar to Fig. 9.29)

Least squares fit only to the two spectra at top and middle  
 (the other dashed curves are computed from the tornado model parameters obtained from the fitting)

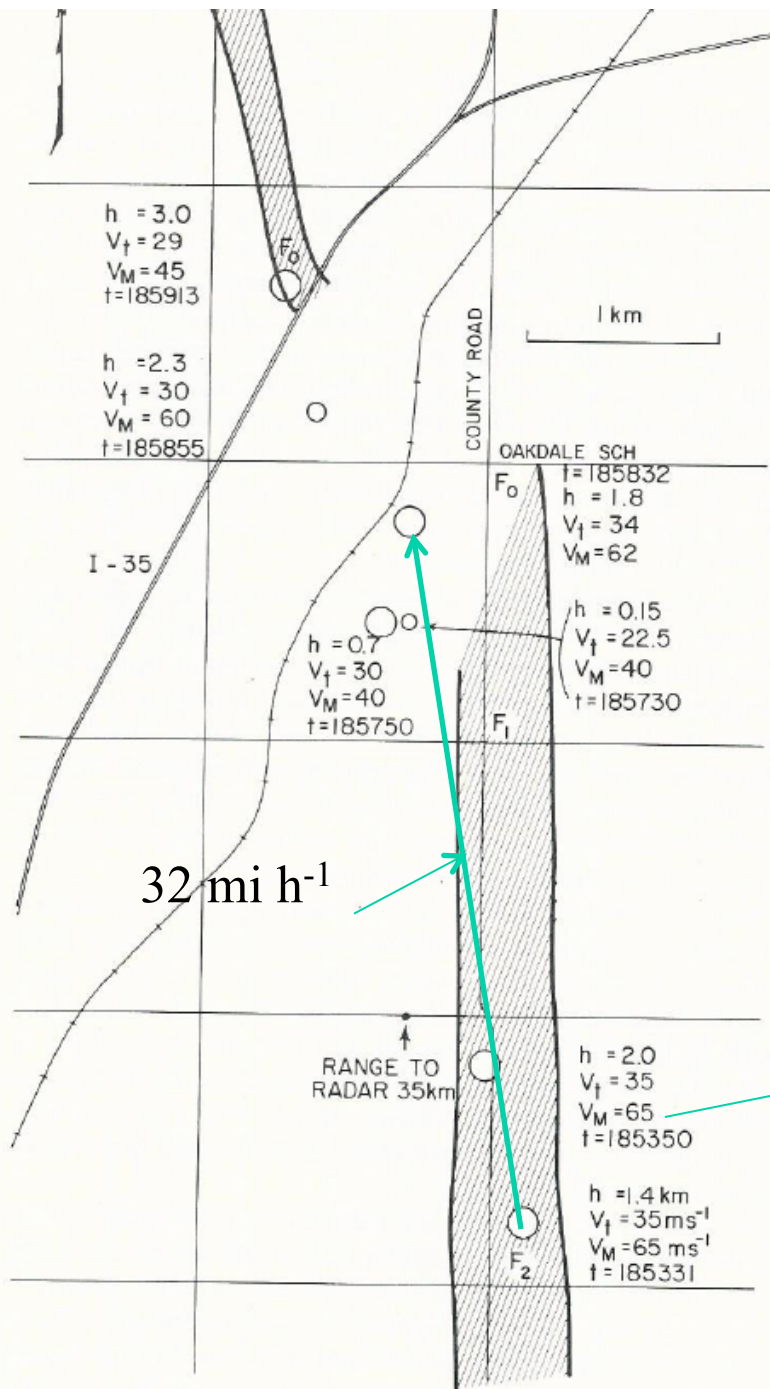
————— Measured spectrum

----- (the other dashed curves are computed from the tornado model parameters obtained from the fitting)



# Fig. 9.30 Del City Tornado (May 20, 1977)

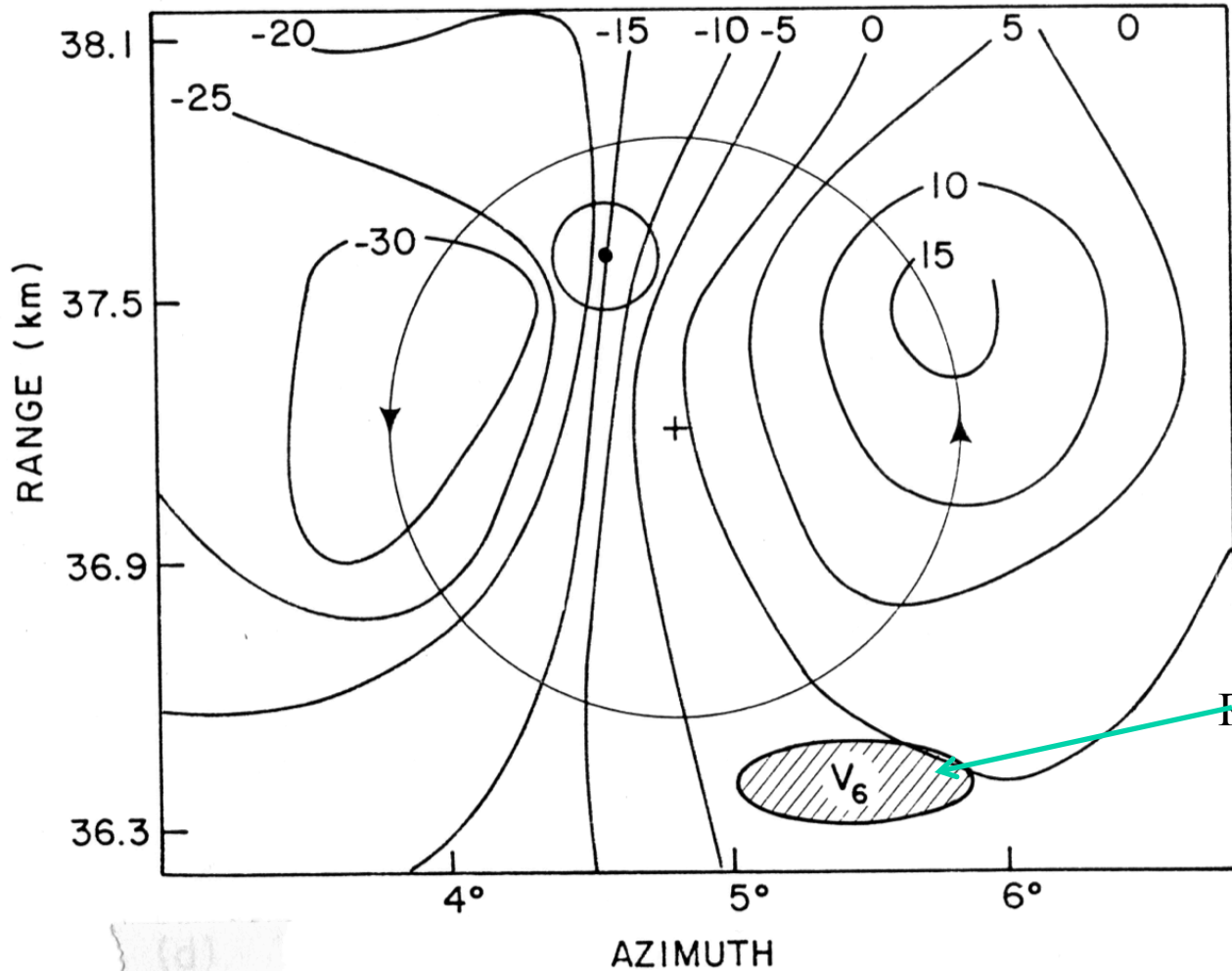
(Tornado Parameters  
Deduced by Fitting  
Doppler Spectra--hatched  
tornado path obtained from  
damage surveys)



# Isodops of Del City Tornado Cyclone (Fig. 9.25)

185748 - 185752 CST

EL = 1.0 ° h = .7 km



Mean Doppler  
data spaced:

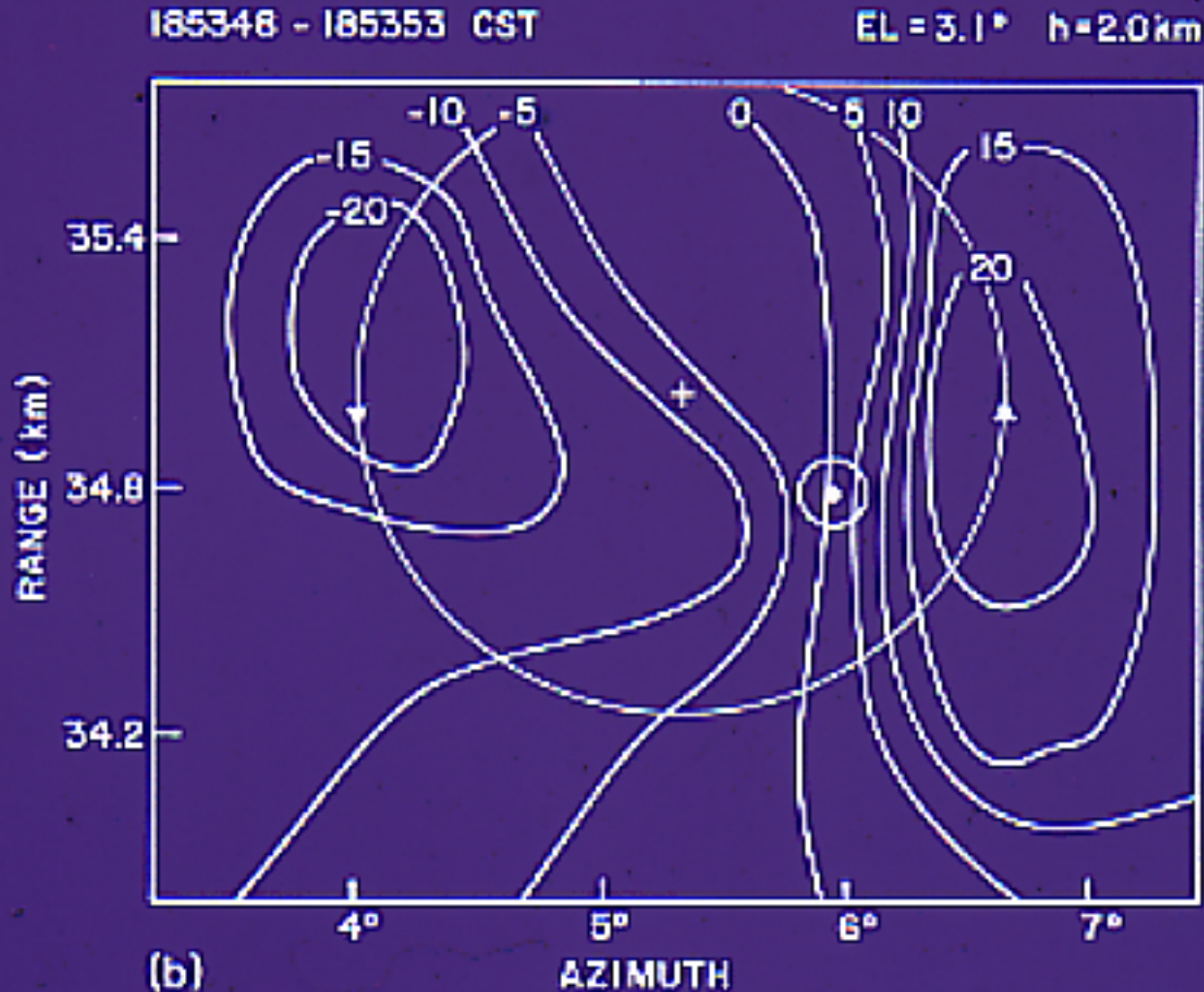
$$\Delta az = 0.2^\circ$$

$$\Delta r = 600 \text{ m}$$

Radar resolution volume

(d)

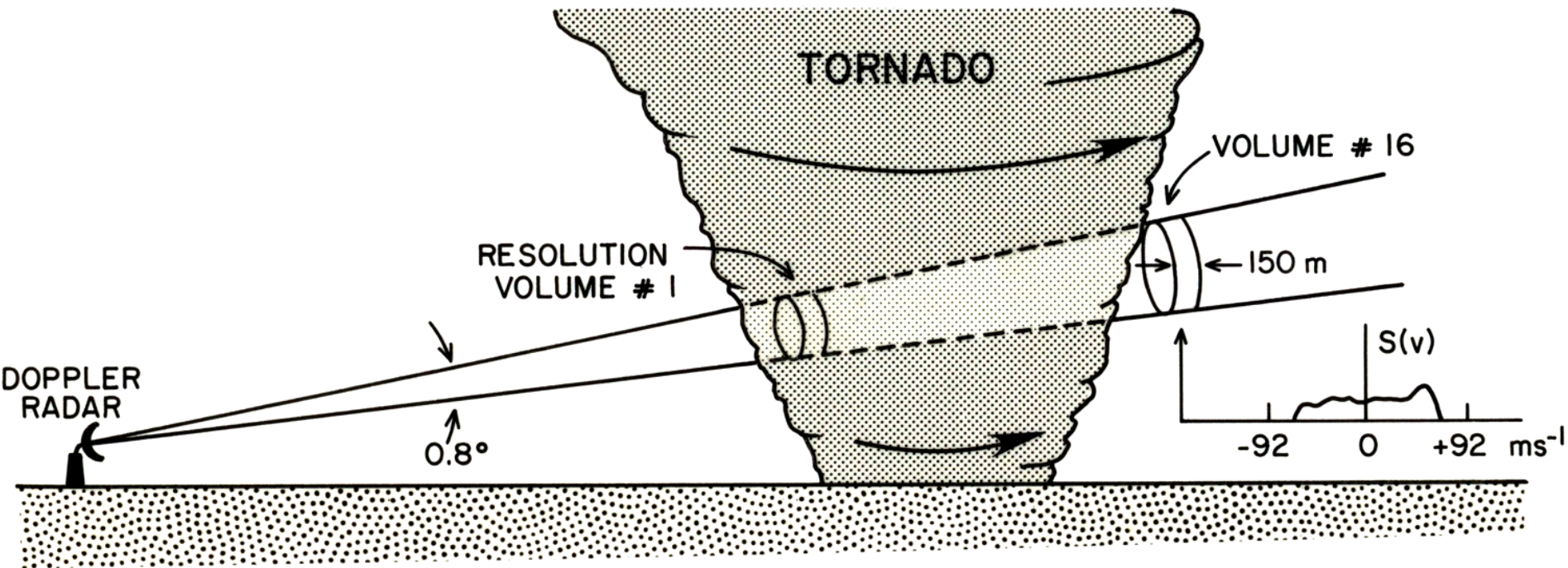
# Isodops of Del City Tornado Cyclone (4 minutes earlier)





# Radar Beam Penetrating a Tornado

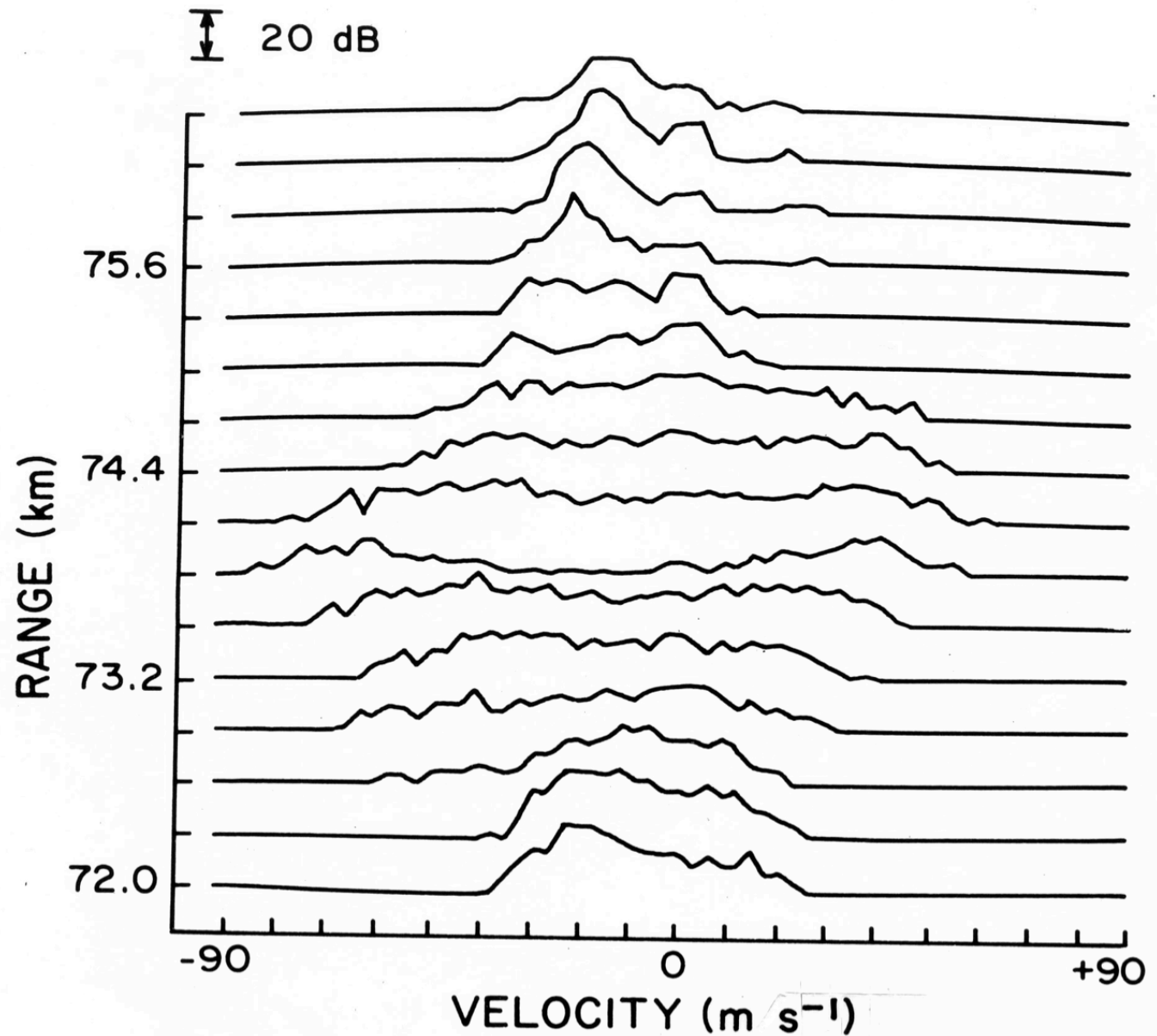
(Fig. 9.28)



# Effective Beam Cross Section and the Binger Tornado (Fig. 9.31a)



# Binger Tornado Spectra (Fig. 9.31b)



**Signal Processing to  
obtain accurate measurements of  
Doppler spectral moments  
and Polarimetric Variables  
(Chapter 6)**

# Goals of Weather Radar Signal Processing

- Extract desired information from received signals
  - **Spectral moments**
    - Reflectivity ( $Z$ )
    - Doppler velocity ( $v$ )
    - Spectrum width ( $\sigma_v$ )
  - **Polarimetric variables**
    - Differential reflectivity ( $Z_{DR}$ )
    - Differential phase ( $\Phi_{DP} \Leftrightarrow K_{DP}$ )
    - Cross-correlation coefficient ( $\rho_{HV}$ )
- For each beam direction there are  $\sim 1000$  range locations probed every  $\sim 1$  ms (lots of data!)
  - Antenna continuously scans the surrounding volume
- **The goal is to obtain the best possible meteorological variable estimates in the shortest possible time (real time)**
  - Remove artifacts
  - Resolve range and velocity ambiguities



U.S. Weather Bureau Forecast Office  
Washington, DC (1926)

Chapter 6 deals mostly with statistical analysis of the variance of the estimates. Although variance of the estimates is important to the use of weather radar data, I have decided to skip any discussion of this topic.

The bottom line of chapter 6:

No matter how accurately we measure **each** weather signal echo sample, there is no way to make perfectly accurate measurements from a single echo sample! This is so because weather echoes are random variables and measurements of one echo sample (e.g., for power measurements), or a pair of samples (for Doppler measurements), has practically no meaning. Thus weather radar must process many echo samples and users of radar data must be content with **estimates of the meteorological parameters of interest.**